1	A Novel Hybrid Machine Learning-based Model for Rockfall Source Identification in
2	Presence of Other Landslide Types Using LiDAR and GIS
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11	
12	Abstract
13	Rockfall is a common phenomenon in mountainous and hilly areas worldwide, including
14	Malaysia. Rockfall source identification is a challenging task in rockfall hazard assessment.
15	The difficulty rise when the area of interest has other landslide types with nearly similar
16	controlling factors. Therefore, this research presented and assessed a hybrid model for rockfall
17	source identification based on the best tested stacking ensemble model of random forest (RF),
18	artificial neural network, Naive Bayes (NB) and logistic regression in addition to Gaussian
19	mixture model (GMM) using high-resolution airborne laser scanning data. GMM was adopted
20	to automatically compute the thresholds of slope angle for various landslide types. Chi-square
21	was utilised to rank and select the conditioning factors for each landslide type. The best fit
22	ensemble model (RF-NB) was then used to produce probability maps, which were used to
23	conduct rockfall source identification in combination with the reclassified slope raster based
24	on the thresholds obtained by the GMM. In the meantime, landslide potential area was

25 structured to reduce the sensitivity and noise of the model to the variations in different 26 conditioning factors for improving its computation performance. The accuracy assessment of

27 the developed model indicates that the model can efficiently identify probable rockfall sources

with receiver operating characteristic curve accuracies of 0.945 and 0.923 on validation and training datasets, respectively. In general, the proposed hybrid model is an effective model for rockfall source identification in the presence of other landslide types with a reasonable generalisation performance.

32 Keywords: Rockfall; Debris flow; Hybrid model; LiDAR; Gaussian Mixture Model

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35 **1 Introduction**

36 Rockfalls are common natural hazard in many places worldwide, including Malaysia with high 37 and steep terrain with presence of discontinuities (Simon et al. 2015). This phenomenon affects 38 transportation ways, communication and urban areas that are situated near steep mountainous 39 and hilly areas. The hazard of rockfall is increasing in mountainous regions due to the growth 40 of population and economic activities (Fanos and Pradhan 2018). Rockfall can be defined as 41 separate boulders released from a cliff with different motion modes: flying, bouncing, rolling, 42 or sliding (Vernes 1984; Pradhan and Fanos 2017a). Such events can cause serious causalities 43 because they are difficult to be predicted and can move rapidly depending on the geometric 44 and geomorphologic characteristics of the moving block.

45 Considerable research has been performed on rockfall hazard around the world including 46 identification of rockfall source areas (Fanos and Pradhan 2016; Losasso et al. 2017), 47 prediction of rockfall trajectories (Pellicani et al. 2016; Fanos et al. 2016), probability 48 assessment (Gigli et al. 2014), analysis of rockfall runout distance (Fanos et al. 2016), 49 evaluation of rockfall bounce height and velocity (Giacomini et al. 2016) and risk analysis 50 (Mitchell and Hungr 2016; Pradhan and Fanos 2017b).

51 In particular, identification of rockfall source areas is required in the assessment of rockfall 52 probability and risk because it controls the trajectory of rockfall. Rockfall sources can be 53 identified through in-situ survey or rockfall inventory dataset. Nevertheless, such techniques 54 are costly, time consuming and require experts in this field who are only few in number. In55 situ and inventory data are also usually unavailable or incomplete in space and time for several 56 regions (Kromer et al. 2017). The availability of geographic information system (GIS) data and 57 accurate 3D surface models has enabled the development of many approaches for rockfall 58 source identification (Loye et al. 2009; Lan et al. 2010; Massey et al. 2014). Existing methods 59 rely on the identification of slope angle threshold angles that are considered unstable. For 60 example, a threshold of >49° was used by Lopez-Saez et al. (Lopez-Saez et al. 2016), whereas 61 $>60^{\circ}$ was utilised in (Corona et al. 2013). Moreover, recently developed approaches rely on 62 slope geometry derived from LiDAR point cloud and other conditioning elements, such as 63 slope, aspect, curvature, block type and landuse, by using statistical, probabilistic and machine 64 learning methods (Guzzetti et al. 2003). In Dickson et al. (2016), identification of unstable 65 rocks was conducted using photogrammetric survey in composite construction regions. Many 66 controlling factors of rockfall movement along slope were assessed by Agliardi et al. (2016). 67 The results showed that rockfall source areas cannot be easily identified because they are 68 controlled by different factors. More recently, Mote et al. (2019) proposed a method for 69 rockfall risk assessment through the characterization of rockfall source areas. They considered 70 the continuous cliff bands with slope steeper than 45° as rockfall source areas. Their result 71 shows that rockfall sources are key element in rockfall risk assessment and designing a 72 mitigation process. However, such method is critical to obtain a realistic result as it is restricted 73 to cliff face and rockfall source areas are controlled by additional conditioning factors.

Landslides probabilities are controlled by various conditioning factors including morphological, hydrological, geological, and anthropogenic factors. However, each factor has different relative significance to landslide probability and considering a big number of conditioning factors could lead to a negative impact on landslide probability modelling thus producing an unrealistic result. On the other hand, structural and geotechnical, such the bedrock setting, the spatial frequency of discontinuities (fractures, cracks, and joints), the spatial

80 orientation of the families of discontinuities also influence the landslide probability mapping.
81 However, such information demands an extensive field geomechanical surveys which are
82 costly and time consuming. In addition, such in-situ surveys are hard to be performed in
83 regional scale study (wide area). This study focus on using LiDAR-based landslide
84 conditioning to examine the performance of laser scanning data for landslide probability as
85 alternative of structural and geotechnical factors.

86 Machine learning techniques, which have become common approaches for modeling landslide 87 susceptibility over large regions. The basic assumption of the empirical approach is that future 88 landslides are likely to occur in similar conditions of the past (Fanos and Pradhan 2016). 89 Algorithms, such as random forest (RF) (Youssef et al. 2016; Chen et al. 2018), artificial neural 90 network (ANN) (Pradhan et al. 2014; Truong et al. 2018), Naive Bayes (NB) (Pradhan et al. 91 2014; Pham et al. 2016) and logistic regression (LR) (Bui et al. 2016; Lombardo and Mai 2018) 92 have been widely employed for landslide probability modelling. On the other hand, ensemble 93 methods have been quite exercised in other fields, nevertheless, the application of these 94 techniques in the assessment of rockfall issues is still rare (Truong et al. 2018). However, the 95 use of ensemble models can improve the result of landslide probability mapping (Evans and 96 Hudak 2007; Chen et al. 2018).

97 Kinta Valley is one of the main districts in Malaysia. The bedrock geology for Kinta Valley 98 and surrounding areas are granitic hills, limestone bedrock, and mine. As a result, a lot of 99 engineering geologic issues have been encountered Kinta Valley and its immediate 100 surroundings, involving rockfalls, debris flow, and shallow landslides. The bedrock of 101 limestone in Kinta Valley rises over the alluvial plains forming limestone hills with vertical to 102 sub- vertical slopes (Simon et al. 2015).

104 The aforementioned studies have made remarkable attempts to propose approaches that can 105 precisely allocate rockfall sources by photogrammetry or with LiDAR data. However, one 106 issue still not considered which is where the analysis area includes other landslide types with 107 nearly the same controlling conditioning factors such as shallow landslide, rockfall and debris 108 flow. Although Fanos et al (2018) tried to identify rockfall source areas using an individual 109 machine learning algorithm. Whereas, ensemble models can produce better accuracy. The 110 optimization of the model hyper-parameters was nor performed. In addition, the slope 111 thresholds were determined based on the inventory dataset not on the morphological units of 112 the slope. Therefore, the current research proposes a hybrid model designed for rockfall source 113 identification based on LiDAR dataset in such conditions (the presence of other landslide 114 types). The proposed model uses three algorithms, namely, Gaussian mixture model (GMM) 115 and stacking random forest (RF) coupled with Naive Bayes (NB) (RF-NB). Kinta Valley 116 encountered several landslide incidents including roclfall, shallow landslide, and debris flow. 117 Thus it was selected to evaluate the proposed hybrid model.

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119 2. The Characteristics of the Study Area

The study area is located at Kinta Valley in the West of Malaysia (Figure 1), which is situated approximately 200 km north of the capital city, Kuala Lumpur. The study area is located approximately between the northeast corner $(101^{\circ}5'30'' \text{ E}, 4^{\circ}34'50'' \text{ N})$ and the southwest corner $(101^{\circ}10'45'' \text{ E}, 4^{\circ}30'40'' \text{ N})$. The study area consists of various landuse features, such as urban, grassland, peat swamp forest, oil palm forest and shrub. The extension of the study area is (5 * 5 km) with landslide density of $(2.28 \text{ event/km}^2)$.

127 The humidity at the study area is relatively high (approximately 82.3%) throughout the year,
128 and the temperature lies between 23 °C and 33 °C (The Meteorological Service Department of
129 Malaysia). The average annual rainfall in Kinta Valley is 323 mm.

130 The geological setting of the Kinta Valley is completely varied with a high percentage of igneous rocks. However, sedimentary (limestone) and metamorphic rocks (marble) are 131 132 profusely present in the district. However, the selected area contains only limestone. Several 133 faults exist in the study area (Pham et al. 2016). limestone hills prone to landslides incidents 134 because of the presence of extensive fractures and joints that can be easily triggered by various 135 factors, such as water saturation. The faults can also increase the potential of landslides 136 occurrences as they triggering earthquakes. Consequently, Kinta Valley has encountered many 137 landslide events including shallow landslide, rockfall and debris flow.



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Fig. 1 Study area, Kinta Valley, Malaysia

141 **3. Materials and Methods**

142 **3.1 The Used Datasets**

The main dataset of this research contained laser scanning data. High-resolution LiDAR point clouds were gathered using an airborne LiDAR system (RIEGL) in 2015 with a flight height of 1000 m. Consequently, high-density point clouds were produced with around 10 pts./m². The collected dataset was processed through GIS to perform filtering and interpolation processes. Processing must be applied to the gathered point clouds to eliminate noises and outliers and produce a precise DTM for extracting the conditioning factors of rockfall.

149 The inventory dataset of landslides is a fundamental element in the assessment of rockfall 150 source areas. This dataset was prepared from different sources including field surveys, remote 151 sensing and historical records. High-resolution aerial photos (0.1 m) that captured during the 152 collection of LiDAR data were utilised for the optical observation of previous landslide events 153 in the study area. Field measurements were also performed using a GNSS system to gather the 154 locations of landslides that occurred underneath vegetated areas or in regions invisible in the 155 aerial photos. This process was conducted using a Global Navigation Satellite System with 156 real-time corrections. Consequently, 87 landslides (28 shallow landslides, 39 rockfall and 20 157 debris flow), as well as their correlated attributes, were obtained for the assessment (Figure 1). 158 The inventory dataset was divided into two groups (training and testing) to assess the accuracy 159 of the proposed hybrid model. Thus, 70% of the inventory dataset was used to build the model, 160 and the remaining data (30%) were used for validation. The dataset was divided into two group 161 randomly insuring the distribution of each group on the whole study area and each group 162 contains all landslide types.

163 **3.2 Deriving of DTM**

164 The collected raw data contained ground and up-ground points. Therefore, a filtering algorithm 165 must be used to eliminate the up-ground points for obtaining an accurate DTM. The LiDAR-166 based DTM should be constructed accurately to extract accurate conditioning factors (Chen et al. 2017). Several approaches have been proposed to perform this process. The current study 167 168 used an algorithm proposed by (Messenzehl et al. 2017) called multi-scale curvature algorithm 169 (MCC) executed within GIS environment. This algorithm can derive an accurate DTM in urban 170 areas with different natural and man-made features (Pham et al. 2017). The terrain details 171 (sharply cut terrains) are essential to rockfall source identification; thus, the window size 172 number should be selected carefully to retain these details (Brenning 2005). Therefore, a 173 particular algorithm was developed to automatically update the number of window sizes for 174 maintaining the details of terrain.

175 The optimal settings of MCC parameters rely on many elements, such as point cloud density, 176 terrain characteristic and the slope interpolation resolution (Chen et al. 2017). Consequently, 177 the MCC parameters of curvature tolerance threshold, scale domain number and convergence 178 threshold were set to 0.3, 3 and 0.1, respectively. After the up-ground points were eliminated 179 through filtering, the inverse weighted distance interpolation technique was used to generate 180 the DTM from the remaining points. Given that the spacing of points was 0.4 m, the DTM was 181 generated with a resolution of 0.5 m. The statistical analysis of the collected point clouds based 182 on root mean square error revealed vertical and horizontal accuracies of 0.15 and 0.3 m, 183 respectively.

184 **3.3 Preparing of Landslide Conditioning Factors**

The source areas of rockfall cannot be assessed on the basis of a certain factor (Agliardi et al. 2016). Thus, the present research used many conditioning factors such as hydrological, morphological, soil and anthropogenic factors to identify the rockfall sources in Kinta Valley.

188 Many factors were extracted from LiDAR dataset, aerial photos and the databases of189 government agencies.

190 Morphological factors (altitude, slope, aspect and curvature) were extracted from the produced 191 0.5 m DTM and GIS spatial analysis tools. The highest altitude in the current research was 375 192 m, whereas the lowest altitude was 37 m (Figure 2a). Slope, which is a major factor that 193 controls rockfall, was utilised (Figure 2b). The aspect ratios were from 0° to 360°, which 194 represent the direction of slope from the north in a clockwise direction (Figure 2c). The second 195 derivative of the DTM was used to calculate the curvature factor (Figure 2d). The curvature 196 controls the flow divergence and convergence across a terrain and the deceleration and 197 acceleration of downslope flows. Therefore, this factor affects deposition and erosion.

198 The topographic roughness index (TRI) is a key hydrological factor that affects landslides199 (Figure 2e). This factor can be calculated using Equation 1:

$$TRI = \sqrt{max^2 - min^2}, \qquad (1)$$

where max is the highest cell value in the nine rectangular neighbourhoods of altitude and minis the minimum value.

203 In the meantime, anthropogenic factors involve landuse/land cover (LULC) and distances to 204 road. Other factors such as distances to stream (derived from a topographic layer) and 205 lineament (derived from an existing map) were also considered in this study. Geological factor 206 is not considered in this research because of the selected study area contains only one type 207 (limestone). Thus, this factor has no impact on landslide probability mapping. In addition, the 208 focus of the current research is on examining the performance of LiDAR deriving landslide 209 conditioning factors. This can increase the generalization of the proposed methods and reduce 210 the model sensitivity to the variation on the conditioning factors. The LULC layer was 211 produced using classified SPOT 5 satellite images with supervised SVM approach 212 (Department of Survey and Mapping Malaysia). Field survey was performed to verify the 213 LULC layer. The landuse map was classified into nine classes: water body, river, 214 transportation, residential building, other buildings, cemetery, forest, mixed vegetation and 215 open land (Figure 2f). Euclidean distance method was used to calculate the distances to road 216 (Figure 2g), river (Figure 2h) and lineament (Figure 2i).

217 Sparsely vegetated areas are more prone to landslide incidents than forests. In the current 218 research, vegetation density was utilised as one of the factors for the rockfall source 219 identification. This factor was derived from SPOT 5 satellite images. Four classes were 220 produced: dense vegetation, moderate vegetation, low vegetation and non-vegetation (Figure 221 2j). Overall, 10 conditioning factors were included in the modelling of rockfall source area 222 identification. Soil texture (Figure 2k) consists of three different types (rocky loam, silt/clay, 223 and loam). This factor is also considered in this research.









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Fig. 2 Landslide conditioning factors

229 **3.4 The Developing of the Proposed Hybrid Model**

This research presents a hybrid model based on two algorithms, namely, ensemble stacking (RF–NB) and GMM, which involved many processing steps, as shown in Figure 3. The major datasets used in this research were landslide inventory map, GIS layers and a DTM derived from airborne LiDAR point clouds. The landslide inventory dataset was utilised to train various ensemble machine learning models and validate the hybrid model. GIS layers including LULC, vegetation density, soil texture, lineament, river and road were adopted to obtain the remaining conditioning factors. The high-resolution DTM was produced using LiDAR point clouds forextracting many factors such as slope, aspect, altitude curvature and TRI.

238 The first processing step is to determine the slope angle threshold of each landslide type 239 automatically based on slope geomorphological units. GMM was run using the slope data that 240 derived from the generated DTM to identify these thresholds. The second step is to determine 241 the best conditioning factors that can identify variance landslide types, including rockfall. This 242 process is performed using Chi-square model as a factor optimisation approach. Consequently, 243 the relevant factors of each landslide type are determined. This process aims to reduce the 244 number of factors for decreasing the time of computation and improving the generalisation 245 capability of the proposed model. The use of only the best factors enables to improve the 246 performance by eliminating redundant and noise information. Thereafter, stacking (RF–NB) 247 model is trained with the inventory data and the selected factors. The stacking (RF–NB) model 248 predicts the landslide probabilities in consideration of the landslide types in the study area. On 249 the other hand, the landslide potential area was constructed. Consequently, a binary raster is 250 generated to reflect the regions that are probable (class 1) and not probable (class 2) to 251 encounter rockfall. This raster is produced through integrating two reclassified elements: slope 252 and landuse. Considering that the study area has encountered many landslides types, the 253 thresholds of slope angle obtained through GMM are used to reclassify the slope raster. The 254 slope raster is reclassified accordingly after the thresholds are estimated automatically. In the 255 meantime, the landuse raster is classified into two classes by integrating water bodies, stream, 256 cemetery, residential building, transportation and other buildings in one class, and the other 257 class contains the remaining classes (forest, vegetated area and open land). The two reclassified 258 elements are integrated to produce the landslide potential area. This process is advantageous 259 because it reduces the sensitivity of the model to the spatial variance in conditioning factors of 260 landslides. In addition, it allows to filter-off the regions with no possibility of landslide. After the thresholds of slope angles are estimated by the GMM method and the likelihood landslide occurrence, the probable source regions can be identified through geoprocessing steps in ArcGIS. Lastly, the remaining data in the inventory dataset are used to validate the obtained results for demonstrating the performance of the proposed ensemble model. The stacking ensemble models were implemented using Python, whereas the GMM was run using Matlab R2016b. The proposed hybrid model was performed in ArcGIS 10.5 environment.



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Fig. 3 Flowchart of the proposed hybrid model

269 **3.5 Determination of slope thresholds**

The distribution of slope angle can be represented in many Gaussian distributions that can reflect the morphological characteristics, such as rock cliff, steep slope, moderate steep, foot slope and plain. A slope is rated as a probable rockfall source area where the slope angle lies over a particular threshold of slope angle, which can be defined through the Gaussian distribution of the morphological unit (rock cliff becomes predominant over the steep slope). GMM comprises k multivariate components normally used as a parametric model for the distributions of landslide probability given by the following equation (Tien Bui et al. 2018):

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$$p(x|\lambda) = \sum_{i=1}^{k} w_i g(x|\mu_i, \sum_i), \quad (2)$$

where *x* is d-dimensional features, w_i , i = 1, ..., k, are the mixture weights and $g(x|\mu_i, \sum_i)$, i = 1, ..., k, are the component Gaussian densities. Each component density is a d-variate Gaussian function of the form

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$$g(x|\mu_i, \sum_i) = \frac{1}{(2\pi)^{\frac{D}{2}} |\sum_i|^{\frac{1}{2}}} exp\left\{-\frac{1}{2}(x-\mu_i)'\sum_i^{-1}(x-\mu_i)\right\},$$
(3)

with mean vector μ_i and covariance matrix \sum_i . The mixture weights satisfy the constraint that $\sum_{i=1}^{k} w_i = 1.$

The GMM parameters were computed on the basis of the training dataset by using the iterativeexpectation-maximisation algorithm.

286 **3.6 Ensemble Machine Learning Models**

287 Machine learning algorithms provide better results for landslide identification than other 288 probabilistic methods. In the last decades, machine learning algorithms have been used 289 effectively in identifying probable landslide areas (Brenning 2005; Evans and Hudak 2007;

Scrucca et al. 2016). Methods, such as RF (Trigila et al. 2003; Chen et al. 2018; Segoni et al. 290 291 2018; Fanos et al. 2018), logistic regression (LR) (Catani et al. 2005; Pradhan et al. 2014; Bui 292 et al. 2016), artificial neural network (ANN) (Manzo et al. 2013; Chen et al. 2017; Pham et al. 293 2017) and NB (Chen et al. 2017, Lombardo and Mai 2018), are popular and widely applied 294 machine learning algorithms for landslide probability and produce high accuracy. However, 295 existing methods for the modelling of landslide probability prove that the forecasting of 296 landslide probability can be improved using hybrid machine learning algorithms (Fanos et al. 297 2018). Thus, new hybrid machine learning models for landslide probability should be 298 developed.

299 The current research partially fills this gap in literature through proposing a new hybrid 300 machine learning model for the probability modelling of different landslide types. Stacking is 301 a machine learning ensemble approach. Contrary to other ensemble models, stacking can create 302 a strong learner from weaker ones with better tuning in the search for landslide probability 303 modelling processes. In comparison with other ensemble models, stacking also requires lesser 304 running time and computational resources for training, optimisation and validation (Alves 305 2017). In this research, different stacking models, namely, (RF-ANN), (RF-NB), (RF-LR), 306 (ANN-NB), (ANN-LR) and (NB-LR), were optimised and trained on the basis of the 307 inventory data and the obtained conditioning factors. The hyperparameters of the used machine 308 algorithms were firstly optimised using the grid search optimisation approach (Kotthoff et al. 309 2017). Then, the best fit stacking ensemble model (RF-NB) was utilised to derive the 310 probability maps of different landslide types. The model was run with 174 samples of the 311 inventory dataset (87 landslides and 87 non-landslides).

312 4 Results and Discussion

313 **4.1 Slope thresholds**

314 The slope angles distribution of various landslide types are presented in Figure 4 based on the 315 inventory data. Various landslide types had occurred at various slope angles, which indicates 316 the potential to identify and recognise the source areas of these types through the GMM. The 317 figure also demonstrates that rockfall incidents had occurred at the highest slope angle range 318 (45–75°). Shallow landslide incidents had occurred within the slope angle in the range from 319 23° to 43° . By contrast, debris flows had occurred at the lowest slope angle range ($15^{\circ}-25^{\circ}$). 320 The thresholds of slope angle depend on the variation in slope angle distribution in a particular 321 region. Thus, the GMM was used to evaluate the ability of determining the thresholds, and the 322 slope angles were fine tuned in an unsupervised way via the GMM algorithm. Consequently, 323 rockfall could be distinguished from other landslide types automatically on the basis of the 324 slope angles.





Fig. 4 Distribution of slope angle for various landslide types in the training dataset

327 The thresholds of slope angles derived via the GMM is illustrated in Figure 5. They included328 five components determined on the basis of the geometric unit of slope terrain. Thresholds

329 were calculated without the label (landslide type). In other words, it is unsupervised process. 330 The mean values (μ_i) of the five components were obtained as follows: 1.46°, 6.23°, 16.43°, 43.21°, 66.31° and 47.22°. Thereafter, the normal values were defined depending on the μ_i 331 332 values in consideration of the standard deviation and mean values of the dataset. This way 333 could determine the efficient thresholds of slope angles. After the slope angles were plotted 334 against the normal values, the effective thresholds of slope angles could be identified through 335 the intersection of curves (slope terrain type), as illustrated in Figure 5. For example, the 336 efficient threshold for debris flow was specified through intersecting the curves of foot slopes 337 with moderate slopes and moderate slopes with steep slopes. This procedure resulted in an effective slope angle in the range from 9° to 23°. For shallow landslide, the effective slope 338 339 angle threshold was determined by intersecting the curves of moderate slopes with steep slopes 340 and steep slopes with cliffs. Consequently, the effective slope angle threshold ranged from 23° 341 to 57°. By contrast, the efficient threshold of rockfall was identified via intersecting of steep 342 slopes with cliffs and above. Therefore, the final threshold was chosen as $> 57^{\circ}$.







Fig. 5 Effective thresholds of slope angles determined through GMM

345 **4.2 Results of Factors Optimisation**

346 Table 1 shows the estimated ranks of the conditioning factors accounting for the different types of landslides, particularly the key factors ($\alpha < 0.05$) (aspect, slope, curvature, TRI, landuse, 347 348 distance to lineaments, distance to streams, distance to roads and vegetation density). Chi-349 square model accuracies (areas under curve (AUC)) are shown with the best conditioning 350 factors. Regarding rockfall, the best five conditioning factors were observed as slope, TRI and 351 distances to lineament, road and stream. However, vegetation density, curvature and aspect 352 were found less significant for the prediction of the rockfall occurrence probabilities in the 353 study area.

Factor	Shallow	Rockfall	Debris Flow	Overall
Aspect	1	9	4	4
Slope	5	1	8	5
Curvature	8	8	7	8
TRI	6	2	3	2
Landuse	4	6	9	7
Distance to lineaments	9	3	6	3
Distance to streams	2	5	5	9
Distance to roads	3	4	1	1
Vegetation density	7	7	2	6
AUC	0.79	0.94	0.88	0.85

Table 1 Factor ranking by Chi-square

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356 **4.3 Results of Stacking Ensemble Models**

The best conditioning factors were derived for each landslide type in the previous section. Consequently, different stacking ensemble models were developed on the basis of machine learning algorithms (RF, ANN, NB and LR) for the prediction of landslide occurrence probability in the study area. These models were trained with the best conditioning factors and 361 the inventory dataset. The success rate curve (ROC) and the prediction rate curve (PRC) were 362 used to assess the performance of each stacking ensemble model. The best fit stacking 363 ensemble model (RF-NB) was used to derive the probability maps of each landslide type. 364 Figure 6 illustrates the generated probability maps. The probability map is raster with spatial resolution of 0.5 m which is the same resolution of the generated DTM. The probability maps 365 366 reflect that shallow landslides could occur in the east of the area. However, higher probability 367 was observed in the steep terrain than in low-slope regions. Some portions in the south and 368 northwest could experience shallow landslides. Figure 6a shows the highly susceptible regions 369 for shallow landslides, which are marked in red colour. In the meantime, the northwest and 370 northeast regions were predicted as highly prone to rockfall. The regions of steep cliffs with 371 high slopes had high probability to encounter rockfall (Figure 6b). Furthermore, the middle 372 towards eastern portions of the study area had high probability to encounter debris flow, 373 particularly the areas with the low slope angle of $< 23^{\circ}$ (Figure 6c).





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Fig. 6 Probabilities of different landslide types

Thereafter, the slope raster was reclassified using the effective thresholds of slope angles to create the landslide potential area raster. A raster with two classes, namely, high potential and less potential of encountering landslides, was obtained. The raster considered landuse and slope angle. The northeast portion, which has steep slopes, was more prone to landslides than others. In general, 24% of the study area could encounter landslides. The next sections demonstrate the results of the developed model to classify these regions depending on the landslide types and transform the probability raster into source areas by utilising the effective thresholds ofslope angles.

384 **4.4 Results of Accuracy Assessment of the Ensembles Models**

The proposed ensemble model was validated using Receiver Operating Characteristic (ROC) 385 386 and precision recall curve (PRC). ROC and PRC explain the known landslide percentage that 387 lay on the rank of the probability level and show the graph of cumulative frequency (Evans and 388 Hudak; Chen et al. 2018). The ROC was produced using the landslide inventory dataset for 389 training, whereas the PRC was produced using the validation landslide dataset. Moreover, the 390 area under curve (AUC) was adopted to assess the accuracy of the tested ensemble models for 391 producing the landslide probability maps; high accuracy is achieved when the area is large 392 (Pradhan et al. 2010; Hong et al. 2015; Wen et al. 2016; Park et al. 2018).

393 Amongst the tested stacking ensemble models, stacking (RF–NB) was found as a best fit model 394 for producing landslide probabilities (Table 2). The highest ROC was found for rockfall 395 (0.935), followed by that for debris flow (0.881). The highest PRC was obtained for rockfall 396 (0.913), followed by that for debris flow (0.859). The model showed the lowest ROC and PRC 397 of 0.805 and 0.797, respectively, for shallow landslides. In general, the proposed model showed 398 weighted averages of 0.889 and 0.856 for ROC and PRC, respectively. The lowest performance 399 accuracy was obtained from the stacking (NB-LR) model with three landslide types. In 400 addition, the stacking (RF–LR) model also proved to be a good ensemble model for predicting 401 landslide probabilities. However, the proposed stacking (RF-NB) ensemble model could be 402 considered an efficient tool because the accuracy assessment revealed an excellent 403 performance of the proposed model based on the validation and training data. Moreover, the 404 model generalisation was expected to be excellent because the PRC of rockfall was higher than 405 that of ROC accuracy, especially in areas with nearly the same characteristics as the tested 406 area. Nevertheless, the accuracy of model performance is also affected by the number of the 407 landslide inventory samples. A realistic model accuracy and result can be achieved with a big 408 number of inventory samples for training and testing dataset. On the other hand, small number 409 of inventory dataset can lead to unrealistic result even with high accuracy achieved through 410 training process. Therefore, the better accuracy achieved in this study is with rockfall dataset 411 due to the big number of inventory samples in comparison with other landslide types. In 412 addition, the lack of the spatial frequency of discontinuities (fractures, cracks, and joints) did 413 not affect the accuracy of the proposed model as it achieved a high accuracy especially with 414 rockfall.

Stacking Model	Debris Flow		Rockfall	Rockfall		Shallow Landslide	
	ROC	PRC	ROC	PRC	ROC	PRC	
RF–ANN	0.820	0.753	0.809	0.785	0.735	0.713	
RF–NB	0.881	0.859	0.935	0.913	0.805	0.797	
ANN–NB	0.795	0.813	0.754	0.739	0.705	0.689	
RF–LR	0.857	0.839	0.874	0.853	0.743	0.755	
NB–LR	0.703	0.675	0.734	0.715	0.659	0.627	
ANN-LR	0.751	0.719	0.795	0.773	0.685	0.667	

415 **Table 2** Accuracy assessment of the proposed model

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417 **4.5 Identification of Rockfall Sources**

The estimated landslide probabilities could be transformed into the source regions by using the efficient thresholds of slope angle derived through the GMM. Subsequently, the reclassified slope raster based on the obtained threshold (>57°) was intersected with the rockfall probability raster within GIS environment to create the probable rockfall source regions. Figure 7 shows the predicted areas of potential rockfall. These regions had steep cliff with other analysed elements (slope components). The model prediction accuracy could be evaluated by determining locations of the recorded rockfall incidents. Most of the historical rockfall incidents (91 %) were accurately predicted through the developed hybrid model. The model predicted that 3.5% (around 0.55 km²) of the area is susceptible to rockfall. The regions that were predicted to be susceptible to rockfall were also investigated through in-situ survey. Many locations were observed to be sensibly predicted as high potential regions to rockfall. These regions were mainly formed by steep cliff surrounded by vegetated areas (Figure 7).





431 **Fig. 7** Identified rockfall source areas using the proposed ensemble model

432 5 Conclusions

This research developed an ensemble model using two algorithms, namely, GMM and stacking
ensemble model based on RF and NB, to identify rockfall source regions in the presence of
other landslide types (shallow landslide and debris flow). The GMM model was used to

determine the effective thresholds of slope angle for different landslide types and construct the
landslide potential area raster. In the meantime, the best landslide conditioning factors were
selected through the Chi-square method. Various ensemble models were developed on the basis
of different machine learning algorithms (RF, ANN, NB and LR). The best fit ensemble model
(stacking RF–NB) was used to produce the probability maps. The binary slope raster created
through GMM was intersected with the rockfall probability map.

442 The developed ensemble model performed well with training and validation regions chosen at 443 Kinta Valley. The model showed accuracies of 0.935 and 0.913 on training and validation 444 datasets. For shallow landslide and debris flow, the proposed ensemble model provided 445 accuracies of 0.805 and 0.881 on the training dataset and 0.797 and 0.859 on the validation 446 dataset. Overall, the proposed ensemble model showed excellent average accuracy on all the 447 landslide types in the inventory dataset. The model achieved weighted average accuracies of 448 0.889 and 0.856 on the training and validation datasets, respectively. Since the proposed model 449 achieved a good accuracy, it proves that the conditioning factors derived from LiDAR can be 450 used as an alternative of the geomechanical factors, such as discontinuity and fractures.

The major contribution of this study is the development of a hybrid model can predict the probable rockfall source regions accurately in the presence of other landslide types. However, additional assessment can be performed to improve the computing performance and accuracy of the proposed model for predicting a particular landslide type in the existence of other types in complex regions.

456 **Funding**

This research is supported by the Centre for Advanced Modelling and Geospatial Information
Systems (CAMGIS) in the University of Technology Sydney (UTS) under Grants
321740.2232335 and 321740.2232357.

460 Acknowledgments

461 The authors acknowledge and appreciate the provision of airborne laser scanning data orthophoto

462 images from airborne laser scanning data (LiDAR) by the Department of Planning.

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